

REGIONAL EM STUDIES IN THE 80's

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Abstract. The review describes in broad terms the development of regional EM studies during the last five-six years. Large simultaneous magnetometer arrays, broadband and dense profiling with five component instruments, the use of remote reference techniques and in-field data processing have increased both the number and the quality of EM surveys. The increase has been strong all over the world.

An extensive list of references, divided geographically, is presented. Selected examples of regional resistivity-versus-depth curves are shown for Africa, the Baikal region, the Baltic Shield, the Canadian Shield, the Carpathian regions, the Central Andes, Iceland, India, the Juan de Fuca Plate, the Münsterland Basin, the Rio Grande rift, the Scottish Caledonides, the Tasman Sea, and for the United States in general. Because of the influence of tectonic settings and the metamorphic grade of rocks, only qualitative aspects of the results are relevant.

'Classical' array studies, especially in Australia, in the Carpathian regions, in India, in North Germany and in Scotland have been reinterpreted and completed with more accurate 2D modelling and dense MT profiling. In the USA and Canada also new regions have been surveyed extensively. New regional EM work has been conducted extensively on the Baltic Shield and in Central and North Africa, Siberia, China, in the areas around the Caspian and Black Seas and in South America.

The newest studies are supported by or compared with other geophysical data, which also are used in extrapolating for missing EM data density. There are several successful large-scale projects in operation: the European Geotraverse (EGT), the KAPG International Geotraverses and the EMSLAB project (with its first preliminary results). Regional EM studies have been increasingly applied to geothermal, hydrocarbon and mineral prospecting as well as local structural studies, e.g. studies of sites for nuclear waste disposal.

Introduction

Electromagnetic induction studies have undergone a drastic increase in volume during the 1980's. A good demonstration of this fact is, that a computer literature search for this review gave more than 400 entries under the password 'MT surveys during 1983 or later'! It is understandable, that only a small part of this material can be included. In this review no attempt has been made to present a unified and critical summary of all the results, but merely some relevant features of studies in different tectonic and geographical settings are illustrated. The author has tried in his selection to emphasize results, which have been modelled at least 2-dimensionally. Because of the great variation of resistivity with the grade of rock metamorphism, tectonic settings etc., it is not useful to pay attention to individual values of resistivity, but instead the qualitative behaviour or the order of magnitude is emphasized in all the summarizing figures of this review.

A considerable number of review articles and books related to the general aspects of regional electromagnetism have been published during the period 1981–1986. (Ádám, 1980/E*; Alabi 1983/E; Campbell *et al.*, 1983/E; Cox,

* The letters refer to the corresponding section in the list of references.

1981/E; Edwards *et al.*, 1981/E; Fournier, 1980/E; Gough, 1983/E; Hermance, 1983/E; Jones, 1983/E; Keller, 1986a, b/D4; Rokityansky, 1982/F; Vanyan and Cox, 1983/E; Vanayan *et al.*, 1983/F). The review and book titles are found in the reference list in Sections E and F.

Induction work of today does not cover only studies in the classical sense (the crust and upper mantle, and their structure as such, on the scale of tectonic units like shields), but also various forms of inductive research techniques have been adapted for a great variety of applications. The applications often cover smaller areas, massifs, intrusions, fault zones, geothermal areas etc. with a dense network of stations. Prospecting for natural resources (minerals, oil, geothermics), and studies of sites for nuclear waste disposal are typical applications (Alperovich *et al.*, 1982/F; Strangway, 1984/E; Ward, 1983/E). One of the more exotic regional studies is involved in the test of the Gold theory of mantle-originating hydrocarbons in Precambrian bedrock (Zhang *et al.*, 1987/A1).

Due to voluminous amount of material it has not been possible to summarize in this review in necessary detail all regional studies of the world in the 1980's. It was therefore decided to concentrate on a few selected examples covering all continents and the most interesting regions of the world, to present some qualitative summarizing figures in the form of resistivity — versus — depth profiles and especially to include a (fairly) complete list of references. Besides language difficulties, quite a few studies, especially of the North American applied regional studies, are published in reports, which appear abundantly as references — but are not easily available. The lists of references have been arranged geographically.

The development in instrumentation has been favourable during the period in question. Large simultaneous magnetometer arrays and dense profiling across interesting conductive structures using MT techniques are typical of regional studies today. A great number of field groups have started use of remote reference techniques and in-field data processing (at least for preliminary work). It is also commonplace to measure all 5 field components during MT surveys.

Mainly in practical applications of the MT technique the use of active (or controlled) sources has gained in success. For crustal studies large generators, most notably the MHD ones developed in the U.S.S.R. (e.g. Velikhov *et al.*, 1986/A1), transatlantic telephone cables (Roy *et al.*, 1985/E) and power lines (McCollor *et al.*, 1982, 1983/D1) have been in use. Since commercially available instruments are rather expensive, many university groups have designed their own instrumentation (e.g. Edwards *et al.*, 1985/D1). Lack of standardization has clearly been a problem in projects where several groups perform measurements in the same area (Hutton 1985/A5).

Typically regional induction studies are increasingly using other than electric data during the interpretation stage, although joint inversion in the strict sense of the phrase has rarely been used. The emphasis on geotraverse type of research efforts all over the world has been favourable in this respect. Other geophysical data can also be used successfully in extrapolating scarce EM data (Hermance *et al.*,

1986/D4). Also attempts to discuss and to understand in detail the tectonic and geological significance of conductive structures and anomalies have greatly increased the value of induction studies for all Earth scientists. The possibility of correlating resistivity models with data from deep boreholes (Büchter, 1984/A3; Jödicke, 1984/A3, 1985/A3) and studies in tectonically active regions like Iceland (cf. Reference Section D2) have been most rewarding in this respect.

Projects, where EM studies are involved, exist on all continents, most of them being related somehow to the International Lithosphere Program (cf. Reference Section A6). Some results from the EGT and KAPG International Geotraverse work are included in the examples of this review, whereas results from the largest EM effort, the EMSLAB project, will be discussed elsewhere.

The emphasis of regional EM studies in the 80's can be said to be twofold: Firstly there has been reconsideration of 'classical' magnetometer array studies by means of reinterpretation, more accurate modelling and dense MT profiling across interesting conductivity anomalies. Such aspects can be clearly seen for studies in Australia (+ Japan and the Pacific, cf. Reference Section A6), in the Carpathians and its surroundings (cf. Reference Section A2), in India (cf. Reference Section B1), in North Germany (cf. Reference Section A3), in Scotland (cf. Reference Section A4), in the U.S.A. (cf. Reference Section D4) and in Canada (cf. Reference Section D1).

Secondly new areas have been covered both by array studies and detailed structural mapping of various kinds. The greatest activity seems here to have been concentrated in the Baltic Shield (cf. Reference Section A1), in Central and North Africa (cf. Reference Section C), in South America (cf. Reference Section D3), in China (cf. Reference Section B3), in Siberia (cf. Reference Section B2), as well as in southeastern parts of Europe and in southwestern parts of Asia.

The improvements in numerical modelling techniques, although not covered in this review, are of the greatest importance in regional studies. Advances have been reported in thin-sheet modelling (Kaikkonen, 1983/A1, 1986/E) and some promising formal interpretation strategies have been presented (a.o. Berdichevsky and Zhdanov, 1985/E; Zhdanov and Frenkel, 1983/E). Perhaps the majority of regional studies already include 2D models of the most important structural features, whereas good, practical 3D modelling is still problematic.

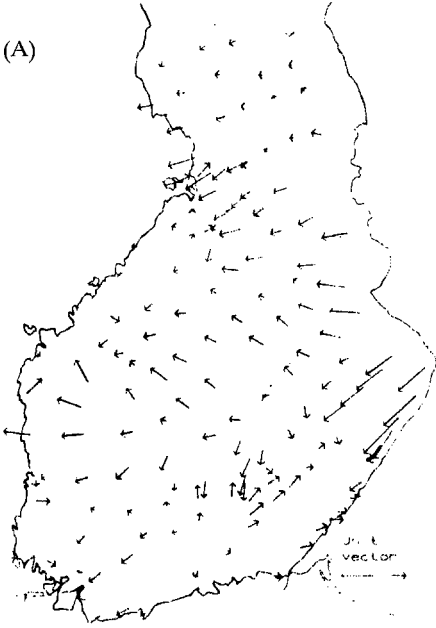
Europe

THE BALTIC SHIELD

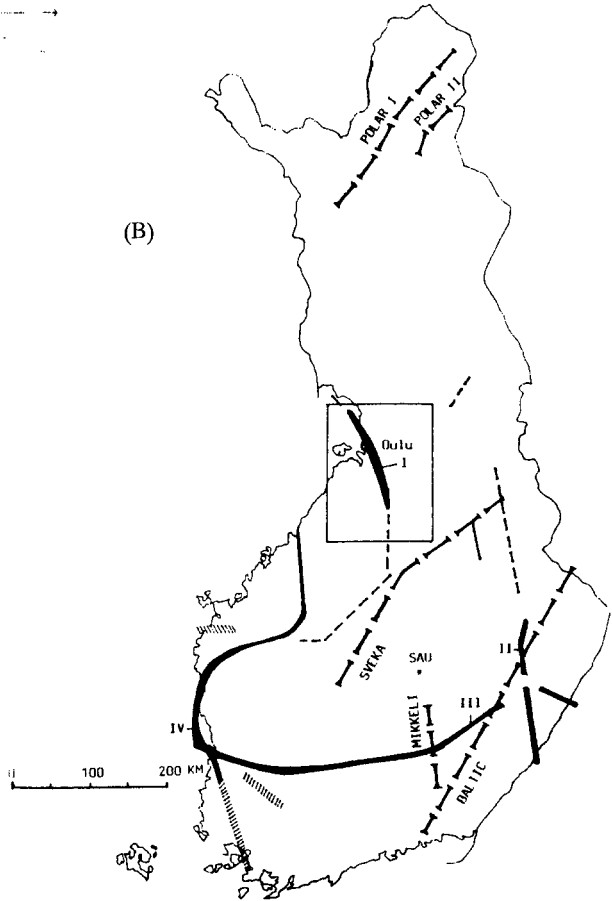
One of the most systematic regional studies of the 80's has been reported from the Baltic Shield, where intense MV array studies have been performed since 1981. The measurements are continuing on the western (Swedish) part of the shield (cf. Pajunpää, 1983/A1, 1984/A1, 1986/A1, Pajunpää *et al.*, 1983/A1; Hjelt *et al.*,

PERIOD 316. SECONDS

(A)



(B)



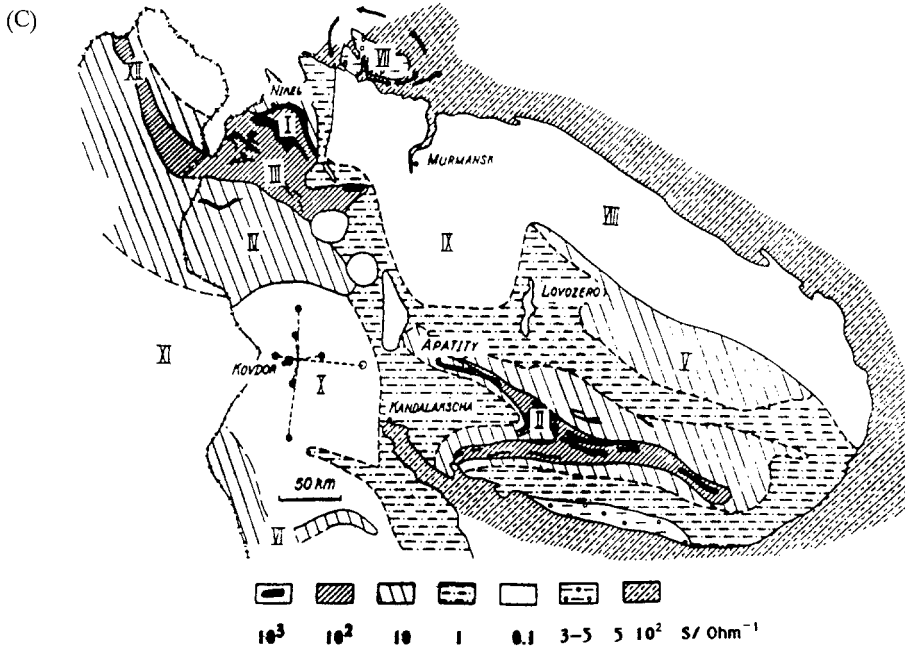


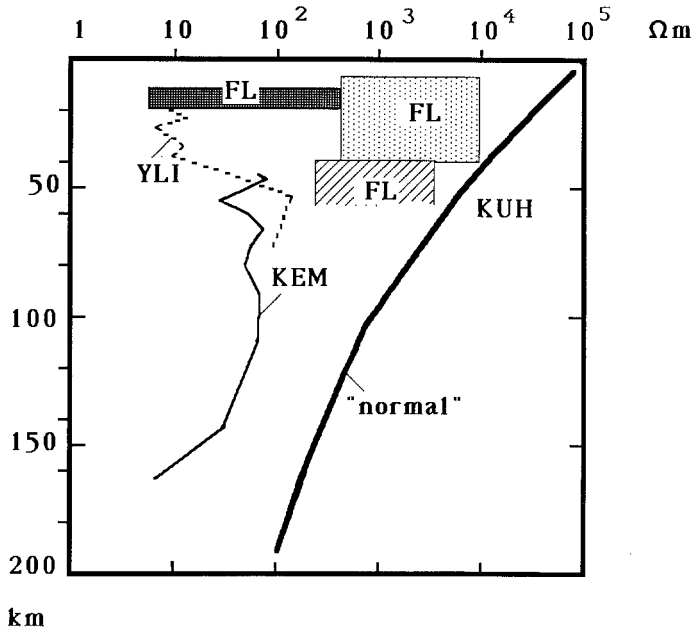
Fig. 1. Examples of the results of induction studies in the Baltic Shield. (A) Reversed induction arrows at $T = 316$ s (Pajunpää 1986/A1). (B) MV conductivity anomalies (Hjelt 1987/A1, adapted from Pajunpää, 1987, in preparation). (C) The S-map for first 10 km of the crust obtained by magnetohydrodynamic soundings (MHDS) (Velikhov *et al.*, 1986/A1).

1986/A1). The array studies found several long bands of good crustal conductors (Figure 1). The upper surface of the conductors starts somewhere between 6 and 12 km and they are several tens of kilometers wide. Complementary MT profiles across relevant crustal anomaly regions are under interpretation (Korja, 1983/A1; Korje *et al.*, 1986/A1; Zhang *et al.*, 1983/A1). Of these the Oulu anomaly has been rather well modelled and seems to require an exceptionally good conductor ($0.5 \Omega\text{-m}$) from a depth of 6 km (cf. Figure 2). Controlled source studies have been undertaken using the MHD generator on the Kola Peninsula (Gorbunov *et al.*, 1979/A1, 1982/A1; Velikhov *et al.*, 1983/A1, 1984a/A1, 1984b/A1, 1986/A1; Heikka, 1983/A1; Heikka *et al.*, 1984/A1). Results of these and earlier studies have been summarized in several papers (Hjelt, 1984/A1, 1987/A1; Hjelt *et al.*, 1986/A1) and in two summarizing reports. A monograph concerned with the Finnish-Soviet joint studies during 1981-1986 is under preparation.

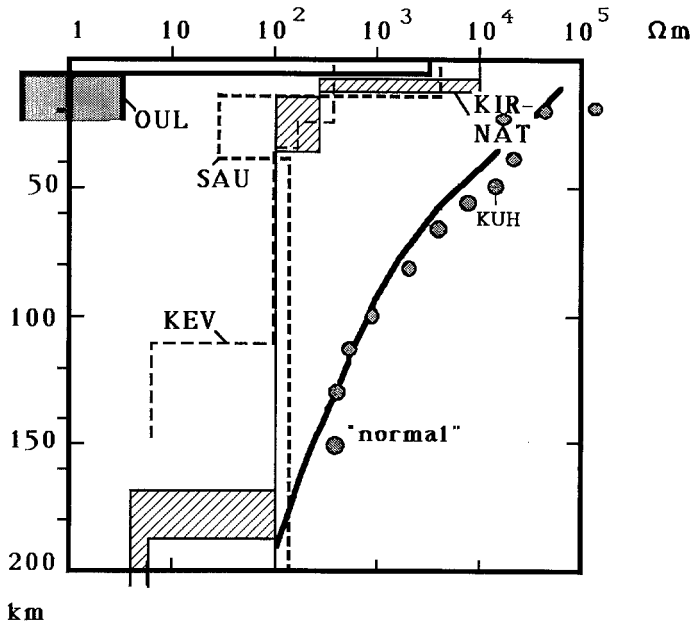
Rather extensive EM studies had previously been performed in the 70's in the Kola Peninsula and in the Soviet Karelia (Zhamaletdinov, 1982/A1, 1984/A1; Golod *et al.*, 1983/A1; Krasnobaeva *et al.*, 1981/A1). The results include S-maps of the Kola Peninsula, the location of the Ladoga crustal anomaly and some proposed regional average resistivity curves (cf. Figure 1C).

Related results exist from the NW part of the Shield, where Jones (1981/A1,

(A)



(B)



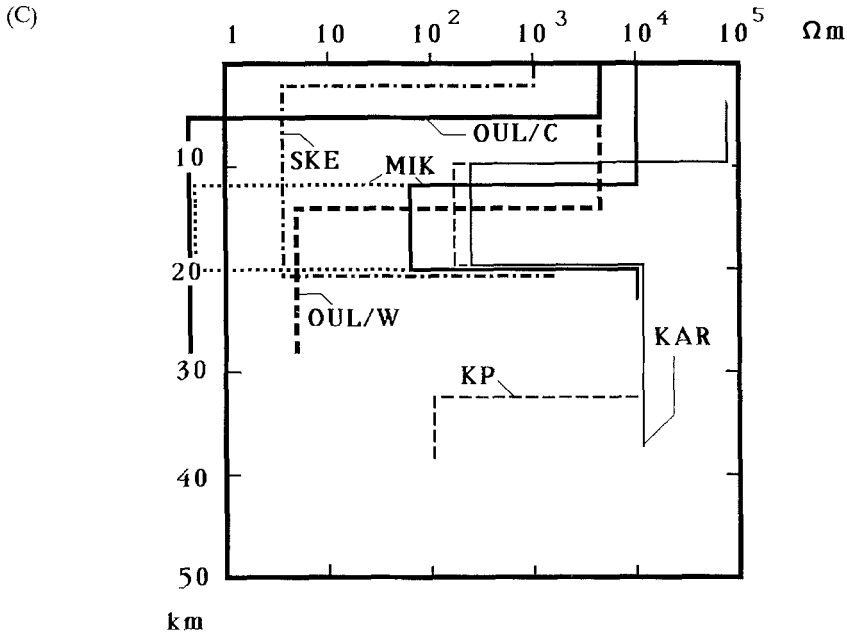


Fig. 2. Qualitative profiles of resistivity versus depth for the Baltic Shield. (A) Selected regional curves, central shield. FL = generalized results (shaded area indicating range of variation) along the Fennolora deep seismic soundings (DSS) profile through E. Sweden (Rasmussen *et al.*, 1987/A1), KEM = Kemijärvi, N. Finland (adapted from Pajunpää, 1987, in prep.), KUH = Kuhmo E. Finland, data points vary around the averaged 'normal' curve (Kaikkonen *et al.*, 1983/A1), YLI = Ylivieska, C. Finland (adapted from Pajunpää, 1987, in prep.). (B) Regional curves, northern shield obtained by Jones (1981–1983/A1, mainly using the Horizontal Spatial Gradient technique). KEV = Kevo, N. Finland, KIR = Kiruna, N. Sweden, NAT = Nattavaara, N. Sweden, SAU = Sauvamäki, Central Finland, For comparison: OUL = Oulu anomaly cross section. (C) Selected conductivity anomalies. KAR = C. Soviet Karelia (Vagin *et al.*, 1985/A1), KP = Kola Peninsula (Vagin *et al.*, 1985/A1), MIK = Mikkeli anomaly, Central Finland (Pajunpää, 1986/A1), OUL = Oulu anomaly N. Central Finland (Korja *et al.*, 1986/A1; Zhang *et al.*, 1983/A1), SKE = Skellefte area, N. Central Sweden (Rasmussen *et al.*, 1987/A1).

1982a/A1, 1982b/A1, 1983/A1) and Jones *et al.* (1983/A1) have analyzed in many ways the IMS magnetometer array data from the late 70's and defined a conducting asthenosphere in Northern Scandinavia (Figure 2B). Osipova *et al.* (1987/A1) have recently pointed out the possibility of an alternative explanation of these data. The source effect problem has gained considerable attention both through direct modelling to set up the limits of a reasonably realistic ionospheric current arch and by using controlled source techniques. Korja (1986b/A1) has used the semipermanent (EISCAT) magnetometer chain for qualitative control of the plane wave conditions of the MTS interpretation along the EGT Polar profile. The region around Kuhmo in eastern Finland gives sounding results most closely related to the undistorted 'normal' resistivity- versus -depth curve for the Shield. (Figure 2A and 2B)

In connection with the EGT studies in Scandinavia a long MT traverse along the FENNOLORA DSS (deep seismic sounding) profile in Sweden and further detailed studies across interesting anomaly regions have been undertaken by the Uppsala group (Rasmussen, 1987/A1; Rasmussen *et al.*, 1983/A1, 1987/A1; Roberts *et al.*, 1983/A1, 1985/A1). Since the distance between the sounding points along FENNOLORA is rather large (40–60 km), only preliminary average values can be given (cf. Figure 2).

NORTHERN GERMANY

EM studies in Germany have a long and famous tradition. In the context of the EGT project the results will be summarized (Berkthold *et al.*, 1984/A3; Berkthold, 1986/A6; Haak, 1985a/A6, 1985b/A6) in unified presentation. Some new information has already been obtained, when Jödicke (1985/A3) and his coworkers were able to identify two separate crustal conducting layers along a NNW-SSE profile in the Münsterland Basin starting from Köln (Figure 3).

Comparison of EM data with deep borehole results (Büchter, 1984/A3; Jödicke, 1984/A3, 1985/A3) together with microscopic studies of the borehole samples has shown the importance of mineral texture and fabric plus submicroscopic film connection between particles for the explanation of the low resistivity layers.

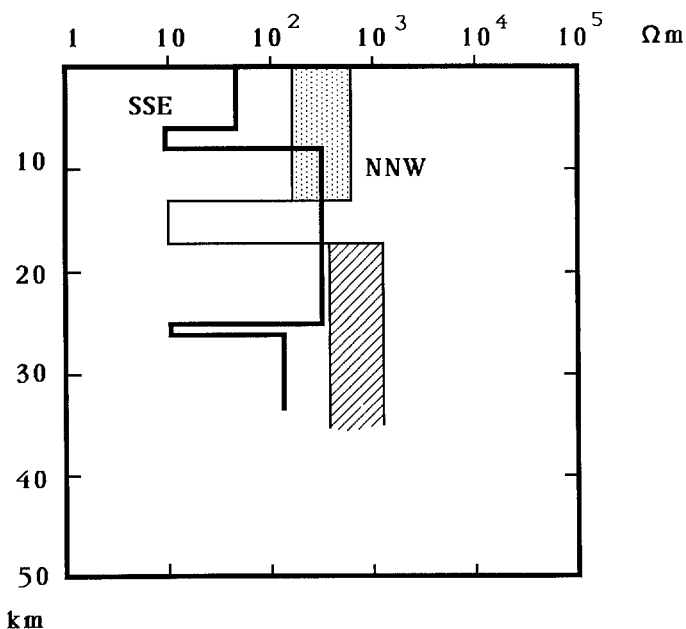
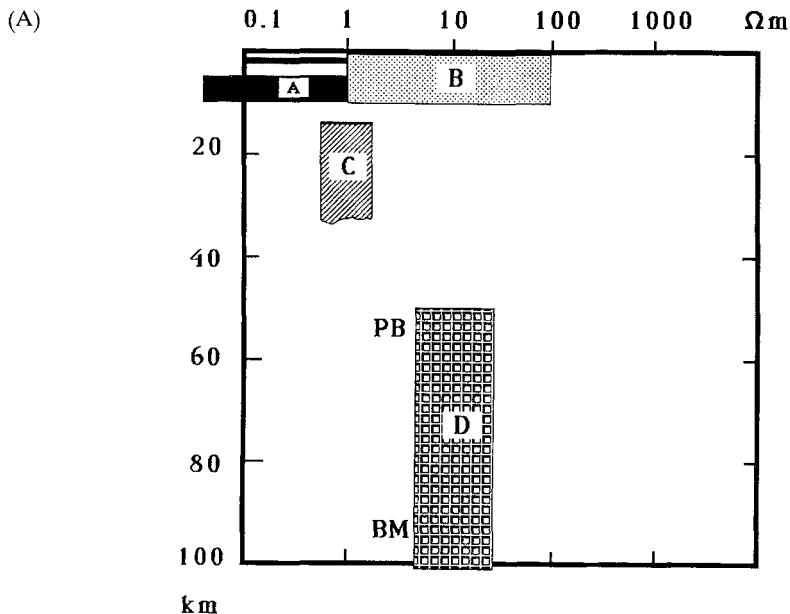


Fig. 3. Selected resistivity-versus-depth curves above the Münsterland Basin after Jödicke *et al.*, (1983/A3). NNW refers to the northern end of the profile, located close to Köln.

THE CARPATHIAN REGION

The Pannonian Basin and the surroundings of the Carpathian arch have been intensively studied targets since the beginning of electrical studies in eastern Central Europe (cf. Figure 4A for a summary of the conducting regions). The understanding of the structure of these tectonically active regions has been approached by scientists from several countries working in close co-operation: Hungary (Ádám, 1981/A2, 1985a/A2, 1985b/A2; Ádám *et al.*, 1983/A2, 1984a/A2, 1984b/A2, 1984c/A2), Poland (Jankowsky *et al.*, 1980/A2, 1985/A2; Bromek, 1982/A2), Czechoslovakia (Praus *et al.*, 1981/A2, 1982/A2; Červ *et al.*, 1984/A2; Pěč *et al.*, 1980/A2; Petr *et al.*, 1984/A2, 1985/A2, 1987/A2; Pichá *et al.*, 1984/A2) and the U.S.S.R. (Rokityansky *et al.*, 1984/A2; Zhdanov *et al.*, 1983/A2, 1984/A2, 1986/A2); recently also Austrian colleagues have participated (Ádám *et al.*, 1981/A2, 1984a/A2, 1986a/A2).

The conductivity structure of the Carpathian arch varies considerably from the E to the W part (Figure 4B). A new crustal anomaly has been reported in the latest results of the Czech MV array studies, but details of the structure have not yet been published. The most interesting findings of the Hungarian studies come from their extension towards the Austrian Alpine region. It has been possible by MTS to follow the Periadriatic lineament, despite topographic and man made noise problems (Ádám *et al.*, 1981/A2, 1984a/A2, 1986a/A2, 1986b/A2).



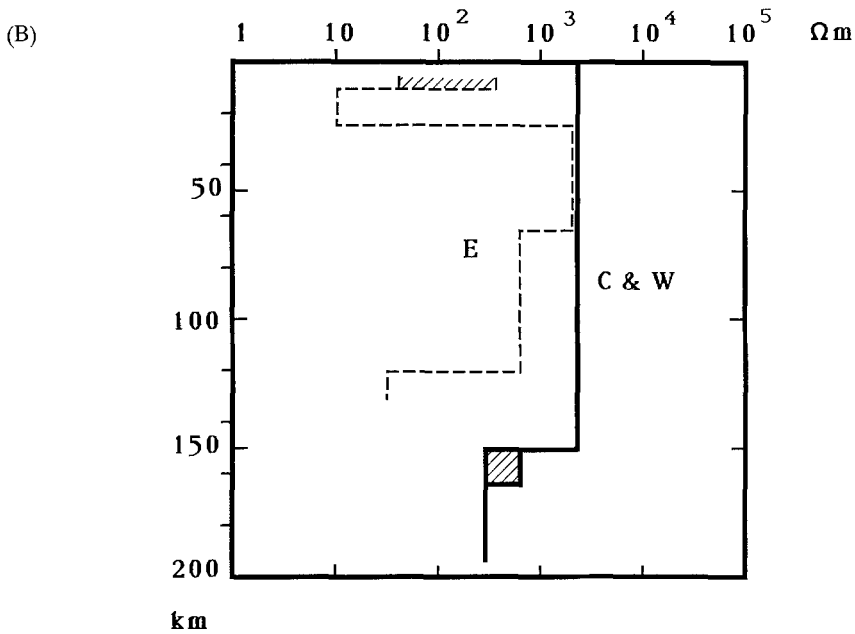


Fig. 4. Qualitative resistivity-versus-depth structures in the Carpathian region. (A) Qualitative indication of conductive zones (after Ádám, 1985b/A2); (A) Graphitic upper crustal conductor; (B) Transdanubian sediments; (C) Carpathian anomaly (melting granite + water; fluids in pores); (D) Mantle conductors (asthenosphere, partially melting periodotite); PB = Pannonian Basin, BM = Bohemian Massiv. (B) Cross-section above the Carpathian arch (from Zhadanov *et al.*, 1986/A2); C, E, and W refer to Central, Eastern and Western parts of the profile across the Soviet Carpathians.

BRITISH ISLES

The regional studies of the British Islands have continued at a regular pace, mostly by scientists of or at least associated with the Edinburgh school of EM studies (for a recent summary, see Hutton *et al.*, 1984/A4). Some selected vertical conductivity sections are presented in Figure 5. The Scottish array studies have been complemented by broadband MTS along selected profiles, the structural models have been improved and algorithms have been developed for 2D inversion. The Edinburgh school has been the main advocate of small site separation in *all* inductive studies. (Hutton *et al.*, 1980/A4, 1981/A4, 1983/A4; Beamish, 1985/A4, Beamish *et al.*, 1983/A4). It is a pleasure to see this approach become more and more accepted, when all realize its benefits in controlling the numerous problems, e.g. surface conductors and 2D/3D structures, involved in any regional (and even more local) induction study.

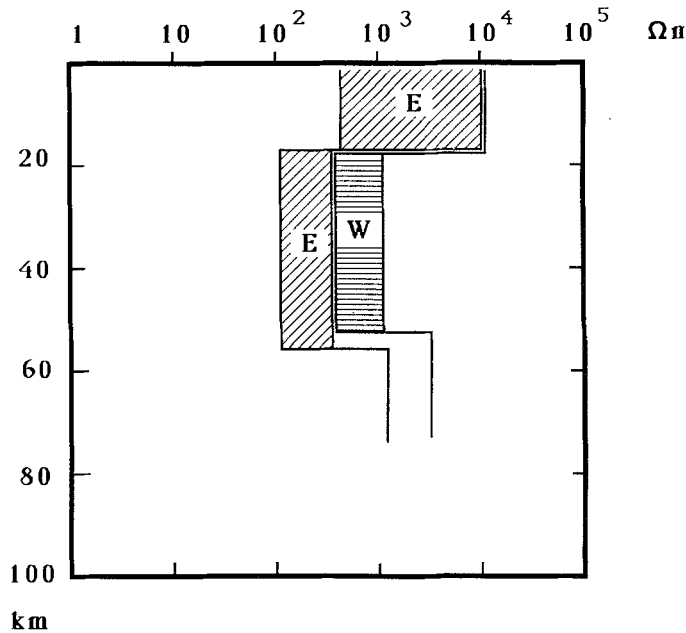


Fig. 5. Qualitative resistivity-versus-depth profiles for the Scottish Caledonides, the metamorphic zone (from Hutton *et al.*, 1983/A4). (E) = region between the Highland Boundary Fault and the Great Glen Fault (W) = region between the Great Glen Fault and the Moine Thrust. The Great Glen Fault itself is modelled as a region of 100–300 Ωm down to the mantle boundary at approximately 50 km.

GEOTRAVERSES AND RELATED STUDIES

EGT

The European Geotraverse (EGT) is a multidisciplinary project, including EM regional studies, trying to establish a crustal cross-section along a profile starting from the African Mediterranean coast to Northern Scandinavia. The profile consists of three segments: Southern, Central and Northern. For the Central section a collection of data for a unified report is proceeding as mentioned above. Unification of instrumentation for long period studies has been proposed by Haak (1985/A6.1).

To the West of the EGT/Central section, but parallel to it, interesting crustal studies have been performed in connection with the French crustal reflection study ECORS. Gazes *et al.* (1986/A6.2) and VanNgoc (1986/A6.2) have reported results from a 15 station MT profile crossing the Paris Basin. The Hercynian crust can electrically be divided into 3 parts, the two northernmost containing conductive layers of 40–60 Ωm lying deeper than 10 km in the North and between 5–10 km in the central part. A still more conducting layer is related to the bottom of the mesozoic basin (5–10 Ωm , $h < 1$ km). A resistivity of 400 Ωm is reported for the bottom part of the crust.

The Northern Segment of the EGT crosses the Baltic Shield and has already been described. MV array work as well as selected MT profiling is continuing in the Swedish part of the Shield.

THE TRAVALE GEOTHERMAL PROJECT

The Southern segment of EGT is parallel to, but West of the Toscana (Tuscany) area, where an intensive, multinational EM study was performed in connection with the European Geothermal Project (cf. the references in Section A5, including the descriptive summary of Hutton, 1985/A5). Although many common electrical features were reported, the problems of nonuniform instrumentation and data processing is evident in many of the reports of this project.

KAPG INTERNATIONAL PROFILES

A similar, but perhaps more systematic approach has been in operation for a long time among the East-European countries in the framework of the KAPG organization. Originally based mainly on DSS studies, the KAPG International Geotraverses increasingly now include also regional EM data. In particular EM results have been reported along KAPG Profiles in the Black Sea and the Caspian Sea areas (Abramova *et al.*, 1984/A2; Ingerova *et al.*, 1986/A6.3) and also in the Carpathians and the North-Poland-region (Stănică *et al.*, 1982/A2, 1984/A2; see also previous references).

Asia, Australia and The Pacific

INDIA

In India several groups have been actively involved in array studies throughout the whole Peninsula. The greatest effort seems to be concentrated in the northern part, the collision zone between the Indian Plate and the Asian plate (Arora *et al.*, 1982/B1, 1987/B1; Lilley *et al.*, 1981/B1; Vozoff, 1984/B1). New details of the conducting structures have been reported, although improvements are still to be expected from subsequent modelling and the completion of the field studies (cf. Curve A, Figure 14).

In the southern part of the Peninsula recent effort seems to have been more on modelling studies than on new field work (Nityananda *et al.*, 1981/B1, 1983/B1; Ramaswamy *et al.*, 1985/B1; Srivastava *et al.*, 1982/B1, 1984/B1; Thakur *et al.*, 1981/B1).

SIBERIA

Studies in the tectonically active Baikal region provide an outstanding example of the benefit of using a statistical and averaging approach in regional studies (Berdi-

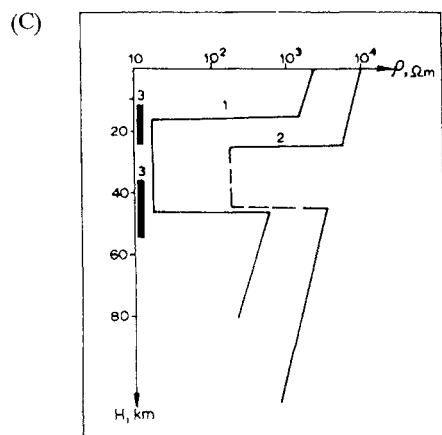
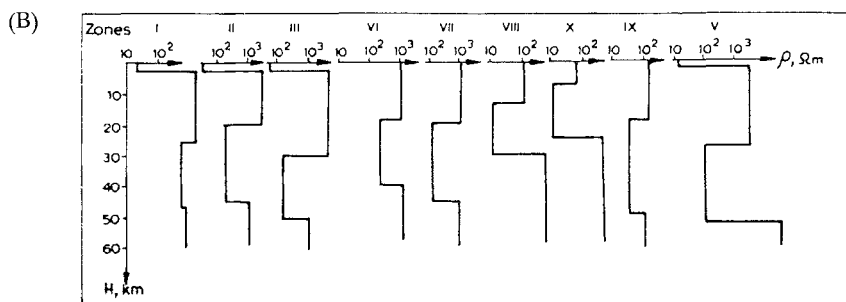
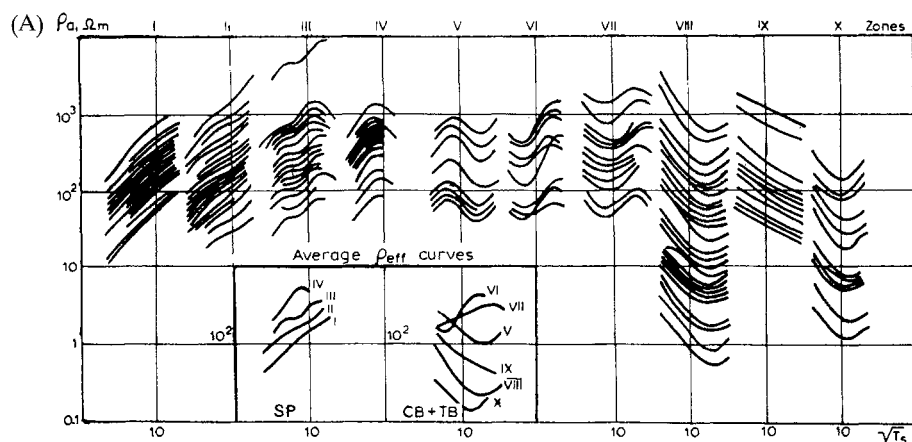


Fig. 6. Statistical regional magnetotelluric (MT) study of the Baikal region. (Berdichevsky *et al.*, 1980). (A) Division of the region into zones using the conformity of ρ_{eff} curves (ρ_{eff} is the geometric mean of ρ_{max} and ρ_{min}). (B) Inversion of the zone-averaged ρ_{eff} curves. (C) Summarized resistivity-versus-depth structures. 1. Trans-Baikal part (TB), 2. Siberian platform (SP), 3. Low velocity layer (LVL) according to DSS.

chevsky *et al.*, 1980/B2; Shilovsky, 1985/B2). Assuming, that the MTS curves are distorted only by galvanic effects (the term static shift seems to be preferred in the English literature) the effective apparent resistivity curves can easily be grouped according to their pattern. Group averages can then form the basis of the final interpretation (Figure 6).

For simplicity references from other Asian regions of the Soviet Union, Sakhalin (Nikiforov *et al.*, 1983/B2), the SW part of Asian U.S.S.R. (Zinger *et al.*, 1984/B2) and Kamchatka (Moroz, 1984/B2) have been included in this section. One of the more recent geoelectrical cross-sections has been constructed across the Amudarja syncline, passing close to the historic city of Samarkand (Berdichevsky *et al.*, 1986/B2).

CHINA

MT activity has been initiated during the 80's in China and Tibet. Several interesting profiles and regional results are reported (cf. the references of Section B3). Reports on the North Margin of the North China Plain seem to contain both crustal and upper mantle conducting layers, although detailed 2D and 3D control still is lacking. From the first few MTS results from Tibet Van Ngoc *et al.* (1983/B3) reports the identification of conductive structures.

JAPAN AND THE PACIFIC

Due to geographical reasons, most results from these areas come from ocean measurements (for an exception, cf. the Honsuy study (Res Group Japan, 1983/B4). Since traditionally oceanic results are reported and reviewed separately at the Induction Workshops, only a list of references for this area is given here without further comments (cf. references in Section B4).

It may be appropriate to add a comment by Hermance (1983/E) concerning sea floor EM studies in general, although it was made in quite another context: 'the most interesting tectonic features of the sea floor are in areas, where all the complications of lateral resistivity variations exist and need to be accounted for'. He continues, that by necessity (equipment limitations and costs) sea-floor data increase more slowly than on land.

AUSTRALIA

EM regional studies in Australia have continued to be actively pursued during the 80's (Lilley *et al.*, 1981a/B5, 1981b/B5, 1981c/B5; Lilley, 1982/B5, 1984/B5). 'Classical' array studies have been continued, details of earlier interpretations have been improved (Chamalaun 1985/B5) and some interesting new structures have been examined first by short profiles (Constable 1985a/B5; Filloux *et al.*, 1985/B5; Finlayson, 1982/B5; Ferguson *et al.*, 1985/B5; Ingham, 1985a/B5, 1985b/B5; White *et al.*, 1985/B5). Constable (1985b/B5) has summarized results from the

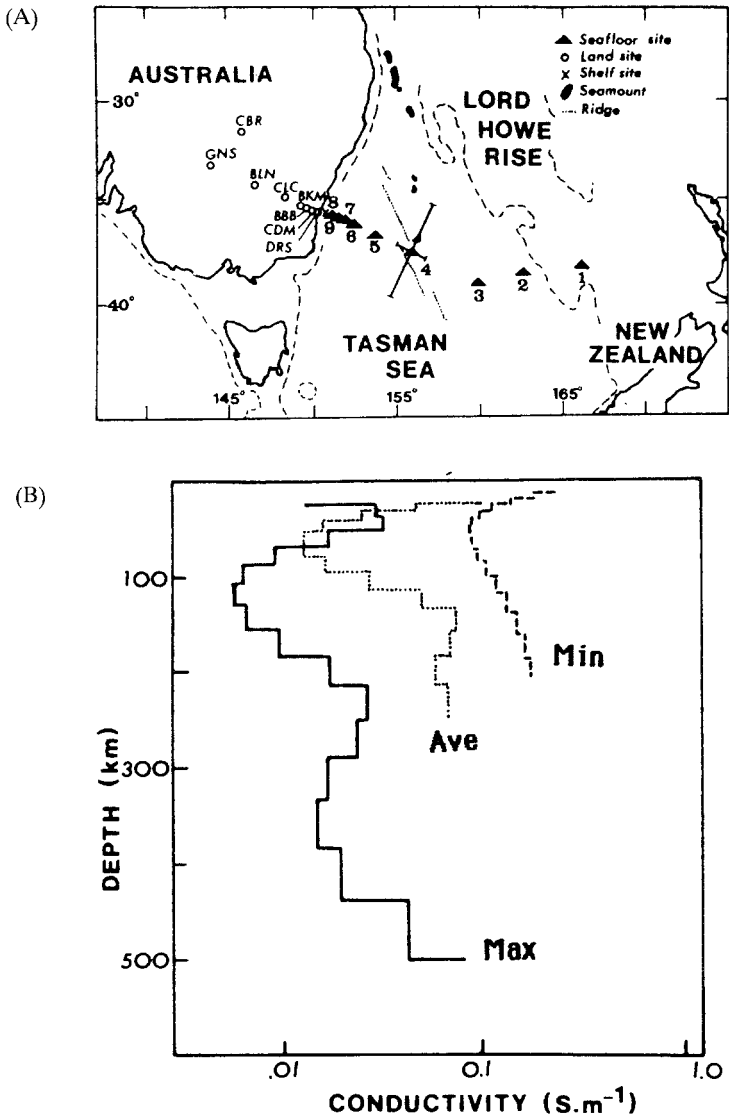


Fig. 7. First results of the Tasman sea floor MT project (Ferguson *et al.*, 1985/B5). (A) Recording sites. (B) Sample 1D inversions of data from site TP4.

Australian continent. Considerable efforts have been necessary to include properly the influence of the ocean on the results. Results from one of the newer profiles, involving sea bottom observations in the Tasman Sea, are seen in Figure 7.

Africa

The 'classical' studies of South Africa have recently been summarized by Gough

(1983/E) and de Beer *et al.* (1982/C). The studies were concentrated on the southern part of the continent, where the mobile Limpopo belt forms a distinctive conductive structure (cf. Figure 8, curve C). More recent regional activities have been concerned with profiles in the middle (Albuoy *et al.* 1982/C; Ogunade, 1983/C; Ritz, 1982a/C, 1982b/C, 1983a/C, 1983b/C, 1984a/C, 1984b/C) and northern (Menvielle *et al.*, 1982/C, 1985/C; Schwarz, 1986/C; Schwartz *et al.*, 1986/C) parts of the continent. Some of the results have been brought together in Figure 8. The Senegal craton seems to be free of a crustal conducting layer, which exists higher up in the crust on the West-African craton than in the mobile belt regions. The crustal conductor appears at comparable depths at the mobile belts both in the Southern and Central parts of the continent. Since the results are based on single profiles only, it seems too early to draw any summarizing conclusions as to the resistivities of the crust on the African continent as a whole.

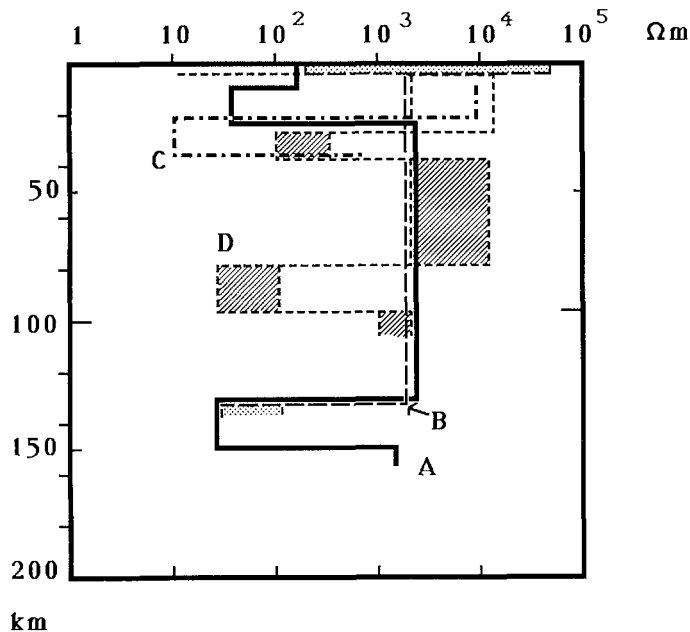


Fig. 8. Qualitative behaviour of resistivity-versus-depth in some African regions. (A) Senegal craton (Ritz, 1982b/C, 1984a, b/C); (B) West-African craton (Ritz, 1982a/C, 1983b/C); (C) The Limpopo mobile belt (Gough, 1983/E); (D) West-African mobile belt (Ritz, 1983a/C).

The American Continents and the Atlantic

CANADA

The activity of Canadian regional electrical studies is on the same high level as in the U.S.A. The most interesting and intense regional effort is the EMSLAB Project in western Canada and U.S.A. and in the surrounding oceanic areas. Since these

results will be reported and summarized elsewhere, only one preliminary result is mentioned here. The upper surface of the subducting Juan de Fuca Plate (cf. curve C of Figure 14) has been located accurately and in good accordance with seismic soundings below Vancouver Island (Gough 1986b/D1; Kurtz *et al.*, 1986/D1).

New details have been added to conductivity studies in 'classical' areas (e.g. Gough *et al.*, 1982/D1; Bingham *et al.*, 1985/D1; Jones *et al.*, 1986c/D1; Hutton *et al.*, 1987/D1), new areas have become the target of electrical regional studies (Jones *et al.*, 1986a/D1, 1986b/D1; Kurtz *et al.*, 1980/D1, 1986/D1; Prugger *et al.*, 1984/D1; Woods, 1985/D1; Woods *et al.*, 1986/D1). Gough (1983/E), Jones (1983/D1) and Kurtz *et al.* (1985/D1) have summarized the geoelectrical results for the Canadian Shield, the main trends being similar to those obtained on the Baltic Shield. The conductive structure of the lower crust, however, is claimed to be more uniform and continuous than is evident from the present Baltic Shield results. The depth range of the Lower Crustal Conductive Layer (LCCL) varies according to geographical position in relation to the Shield boundary and was used as a basis of a classification scheme by Jones (1981/E). Selective profiles from the Canadian Shield are represented in Figure 9, where the Kapuskasing zone is dipping 45° from the surface down to a depth of 10 km and contains at its bottom edge a thin, 10 Ω -m conducting layer. The crustal conductor below Ste Mathilde is much more conductive and the geoelectrical structure below Prince Edward Island has shallow depth conductors, which is assumed to be associated with genuine anisotropy of the rocks (Jones *et al.*, 1986/D1).

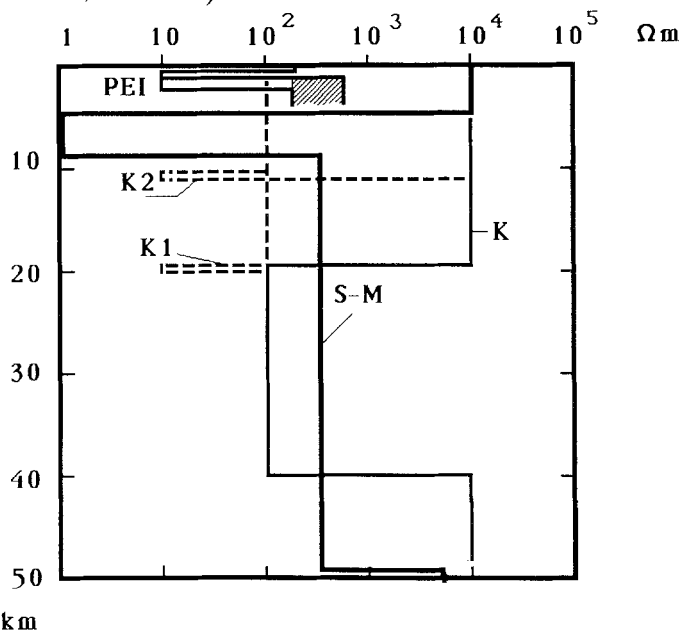


Fig. 9. Selected resistivity-versus-depth profiles for the Canadian Shield. K = the Kapuskasing zone, outside the 45° dipping zone; K1, K2 = two profiles above the 45° dipping zone (from Woods *et al.*, 1986/D1); PEI = Prince Edward Island (from Jones *et al.*, 1986b/D1); S-M = Ste Mathilde (from Kurtz *et al.*, 1980/D1).

ICELAND

Iceland is one of the most fascinating areas for EM crustal research because of its tectonic activity. Several groups have continued their electrical studies on Iceland (Hermance, 1981/D2; Beblo *et al.*, 1983/D2, 1985/D2; Hersir *et al.*, 1984/D2). The results include horizontal resistivity maps at different depths, obtained by DCS (Flovenz *et al.*, 1985/D2), and MTS profiles across the Reykjanes Ridge (Eysteinson *et al.*, 1985/D2; Thayer *et al.*, 1984/D2). The well conducting layers exist at shallower depth above the spreading ridge, whereas both the depth and the conductivity decrease as the distance on either side of the ridge increases (Figure 10, see also curve B in Figure 14).

In connection with interpretation of these regional results, attention has been paid to the mechanisms believed to explain the measured resistivity values. The results of Schmeling (1984a/D2, 1984b/D2, 1985/D2) shed new light upon the problem of the water and melt contents of magmatic rocks as well as on the role of the texture of the minerals grains for determining the resistivities of these rocks.

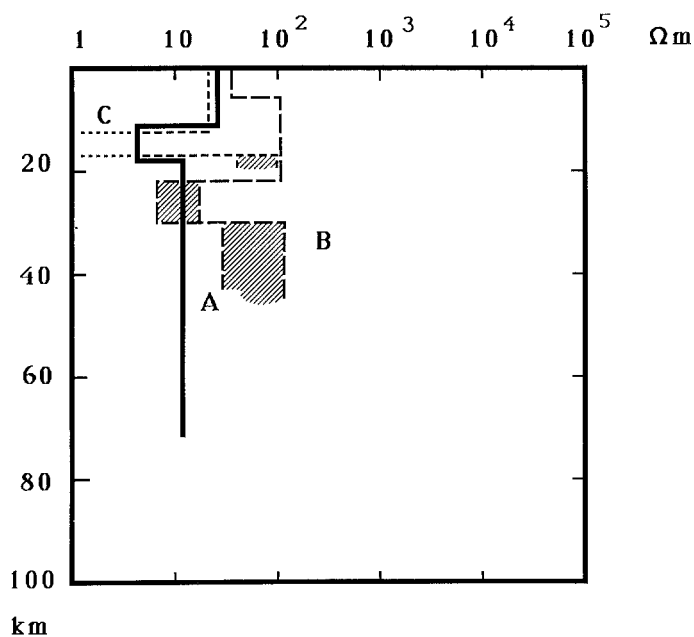


Fig. 10. Selected resistivity-versus-depth profiles for Iceland. (A) Above the Reykjanes rift (Beblo *et al.*, 1985/D2); (B) Far from the Reykjanes rift (Beblo *et al.*, 1985/D2); (C) "Generalized" curve (Eysteinson *et al.*, 1985/D2).

SOUTH AMERICA

The main efforts of regional EM studies in South America have concentrated in two areas. In northern Chile and Bolivia the structure of the Nazca plate and its subducting part beneath the continent has been studied (Haak *et al.*, 1986/D38; Schwarz 1985/D3; Schwarz *et al.*, 1984/D3, 1985a/D3, 1985b/D3). An improved

structural model of the Andean conductivity anomaly has been presented by Tarits *et al.* (1986/D3). The results of Schwarz suggest, that beneath the Altiplano a crustal well conducting structure is indicated above the subducting plate proper (cf. Figure 11). The conductivity increase is interpreted as caused by the upflow of hot

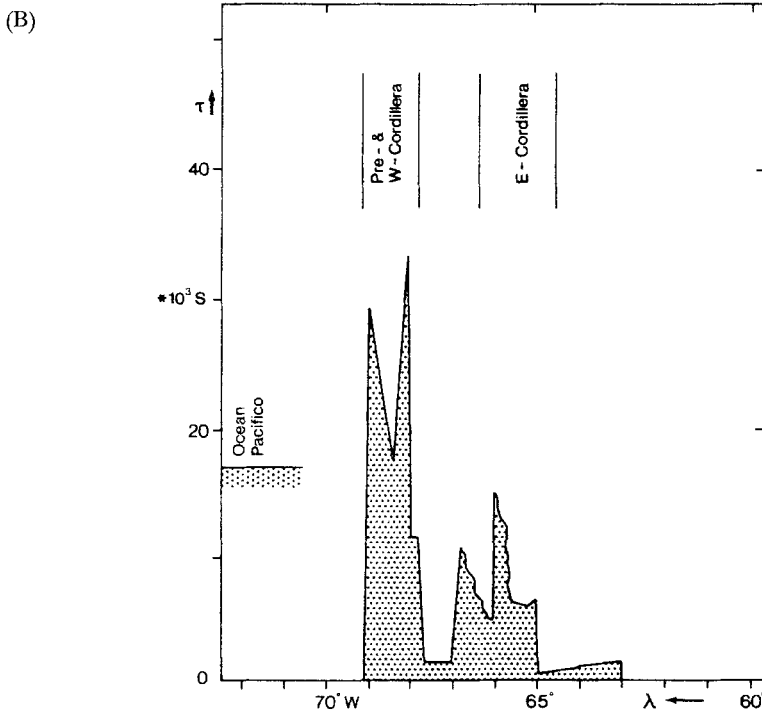
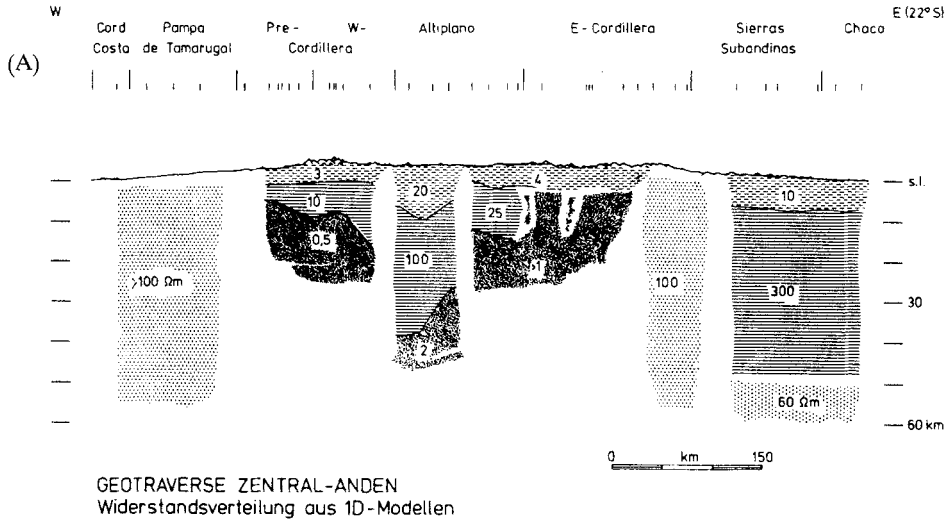


Fig. 11 Conductivity sections across the Central Andes (Schwarz *et al.*, 1985/D3). (A) Geoelectric cross-section obtained by Schmucker-inversion. (B) Integrated conductance for the uppermost 25 km.

material, but the penetration of the measurements is not great enough to indicate the conductivity of the subducting plate itself.

In the southern part of the South American continent EM activities have been reported mainly from Argentina (e.g. Baldis *et al.*, 1983/D3; Febrer *et al.*, 1980/D3; Gasco, 1982/D3; Vatin-Perignon *et al.*, 1985/D3). Many of the results are still at a preliminary interpretation stage.

U.S.A.

Very intensive activity has taken place over the whole North American continent during the last 10 y. The practical application to geothermal and hydrocarbon prospecting has increased the number of soundings tremendously. Keller has reported more than 10000 MTS points in the U.S.A. alone (Keller, 1986/D4, 1987/D4) and presented a concise summary of these results. Only one example of the summarized results is presented here (Figure 12). Keller notes, that although the regional conductivity structures are consistent, they are often at variance with seismic and gravity models from corresponding regions. This shows the need for new data and further detailed studies of existing EM data.

The case history of Jiracek *et al.* (1987/D4) reports on two parallel, regional MT profiles across the Rio Grande rift. The example is an excellent demonstration of the care needed in translating the inverted MTS results to a geological-tectonical model.

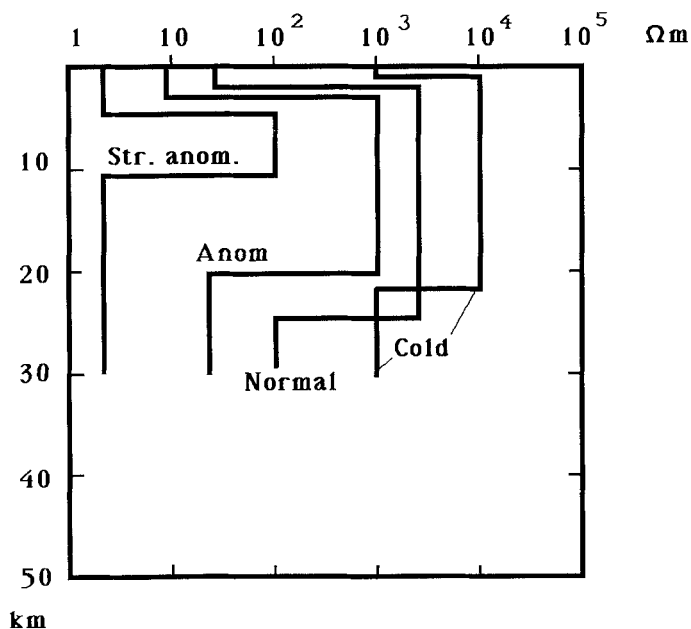


Fig. 12. Qualitative average resistivity-versus-depth profiles for the U.S.A. (from Keller, 1986/D4). *Strongly anomalous* and *anomalous* regions refer to areas in the extensional areas in Western US (Basin and Range province and the Rio Grande Rift); *cold* curve refers to the region around the Great Lakes, thus effectively to the Canadian Shield; *normal* values refer to all the other areas.

It is known, that magmatic activity exists below the centre part of both profiles. However, the conductivity structures beneath the profiles are remarkably different. Jiracek *et al.* (1987/D4) propose an explanation related to the fracturing of the bedrock above the magma chambers. A ductile rock cap is believed to exist below the N profile, trapping the magmatic liquids and forming a well conducting region clearly visible in MTS. In the S profile fluid filled cracks are thought to penetrate the ductile layer allowing the liquids to migrate upwards. No conductive trap is formed (Figure 13). The mountainous parts of the profiles are less conductive throughout.

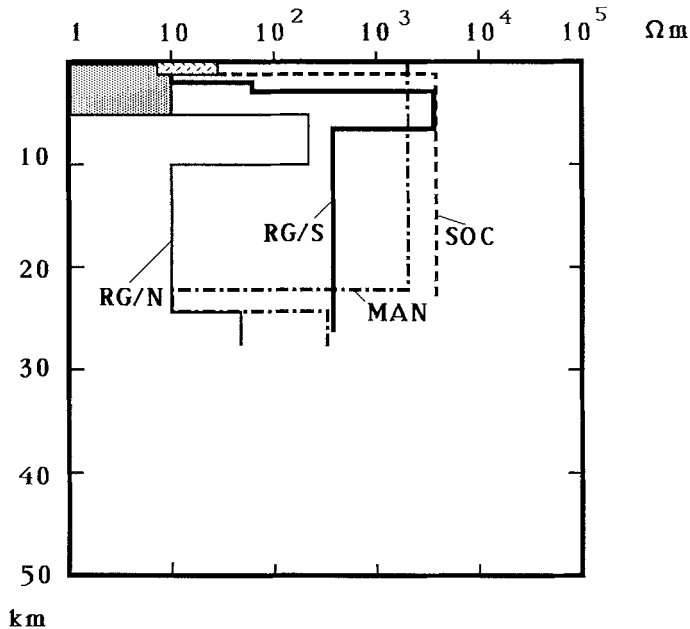


Fig. 13. Selected resistivity-versus-depth profiles for the Rio Grande Rift (from Jiracek *et al.*, 1987/D4). RG stands for locations on the rift itself, S for the Southern line and N for the Northern line; SOC = Socorro Mountains on the S-line, MAN = Manzano Mountains on the N-line.

General Comments

Finally Section E in the list of references mentions some papers of general interest (methodologically or otherwise) as well as some earlier reviews related to regional studies. Quite a few worthwhile books (Section F) have appeared during the 80's, starting with a reprint collection of classical MTS papers (Vozoff, 1986/F) to advanced theoretical monographs on interpretation (Berdichevsky *et al.* 1985/F) and basic theory (Kaufmann and Keller, 1983/F). It is now relatively easy for a research student to obtain a solid background of the methods and interpretation from such texts by himself, whereas only patient, accurate and careful field work forms the basis of improved understanding of the geoelectric structures of the Earth.

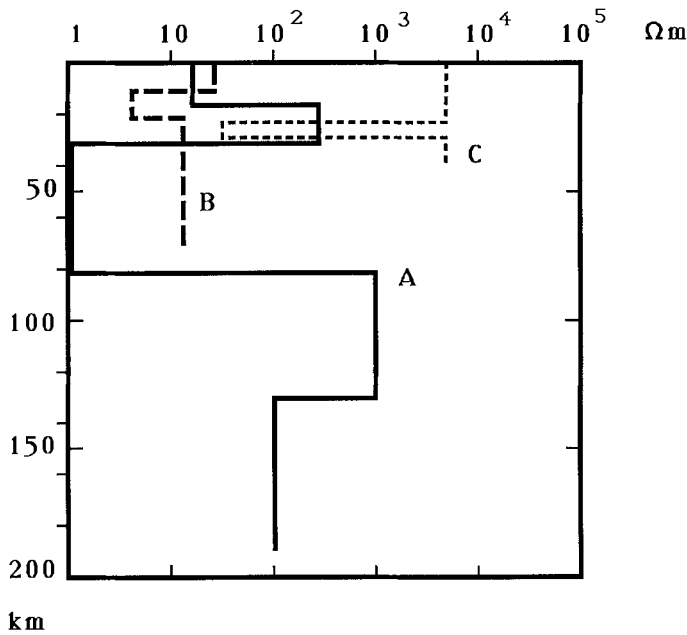


Fig. 14. Selected resistivity-versus-depth profiles for some tectonically active regions. (A) Subduction of the Indian plate under the Himalayas (from manuscript of Chamalaun *et al.*, 1986/B1); (B) The Reykjanes Ridge, Iceland (Beblo *et al.*, 1985/D2); (C) Subduction of the Juan de Fuca plate under the Vancouver Island (Kurtz *et al.*, 1986/D1).

The most outstanding feature when preparing this review has been the tremendous increase in the volume of regional or semi-regional geoelectrical studies all over the world. Since at the same time improvements in instrumentation and especially data processing have been very rapid, it is no longer possible for a single scientist to make a concise and reliable summary of all the results. There is certainly a great need for educated reviews and summaries by region. Hopefully the monograph on the findings of the recent activities on the Baltic Shield and the forthcoming results of the EMSLAB Project will form a good start in this direction.

Summary

A voluminous amount of regional EM studies have been reported during the first half of the 80's. The well-established concepts in instrumentation, data processing and interpretation techniques have made it easier for many new research groups to enter the field of geoelectrical studies. The use of methods earlier regarded only as research tools, have become feasible for practical application, such as prospecting for natural resources (minerals, oil, geothermics), and studies of sites for nuclear waste disposal. The applications often cover smaller areas, massifs, intrusions, fault zones, geothermal areas etc. with a dense network of stations.

Large simultaneous magnetometer arrays and dense profiling across interesting conductive structures using MT techniques are typical of regional studies today. The remote reference techniques, measuring all 5 field components during MT surveys and in-field data processing (at least for preliminary work) has become very popular. The use of active (or controlled) sources have gained in success, mainly in applications. For crustal studies large generators (e.g. MHD), transatlantic telephone cables and power lines have been in use. The lack of standardization of instrumentation and data processing techniques makes comparison over large areas difficult.

Electromagnetic data are increasingly often used together with other geophysical data during the interpretation stage, although joint inversion in the strict sense is still rare. Also discussion and detailed understanding of the tectonic and geological significance of conductive structures is continuously increasing. The possibility of correlating resistivity models with data from deep boreholes and studies in tectonically active regions like Iceland have been most useful. It has become more and more evident, that the accurate knowledge of the resistivity values of various bedrocks, minerals and even regions is usually less important than the overall qualitative behaviour, the order of magnitude and tectonic aspects of the resistivity structures. Statistical averaging techniques are especially valuable in regional MTS surveys, the more so, when sounding bandwidth and station spacing have been inadequate.

Projects, where EM studies are involved, exist on all continents, most of them being somehow related to the International Lithosphere Program. Results from the EGT and KAPG Geotraverse work are included in the examples of this review, whereas results from the largest EM effort, the EMSLAB project, will be discussed elsewhere.

The emphasis of regional EM studies in the 80's can be said to be twofold: Firstly improvement of 'classical' array studies have been made by means of reinterpretation, more accurate modelling and dense MT profiling across conductivity anomalies. Such aspects have dominated the regional studies in Australia, in the Carpathians, in India, in North Germany and in Scotland, but have also taken place in the U.S.A. and Canada.

Secondly new areas have been covered both by array studies and detailed structural mapping of various kinds. The greatest activity has been in the Baltic Shield, in Central and North Africa, in southeastern parts of Europe, in southwestern parts of Asia and in South America.

Recent improvements in numerical modelling techniques, although not covered in this review, are of greatest importance for regional studies. Great advances in thin-sheet modelling and promising new formal interpretation strategies have made it possible, that the majority of regional studies already include 2D models of the most important structural features. On the other hand, good, practical 3D modelling is still not at our disposal.

Active world-wide planning in the context of the ILP is currently aiming to start combined geophysical and geological Geotraverses called TRANSECTS. It is

believed that EM studies can now make a valuable contribution to this programme and all induction scientist should be encouraged to participate in these multimethod geoscientific projects.

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